

# A Finite Dimensional Model of Ice-Albedo Feedback

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## Ice-Albedo Feedback



### Acknowledgements

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## Ice-Albedo Feedback

**Budyko-Sellers Model**

$$R \frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T)$$

insolation      albedo      re-radiation      transport

$T = T(y, t)$ : annual mean surface temperature  
 $y = \sin(\text{latitude}) \quad y \in [0, 1]$   
 $Q$ : global annual mean insolation  
 $s(t)$ : relative annual mean insolation  $\int_0^1 s(y) dy = 1$   
 $y = \eta$ : ice boundary  
 $\alpha(y, \eta) = \begin{cases} \alpha_1, & y < \eta, \text{ albedo} \\ \alpha_2, & y > \eta. \end{cases}$   
 $\bar{T}(t) = \int_0^1 T(y, t) dy$ : global annual mean temperature

## Ice-Albedo Feedback

**Budyko-Sellers Model**

$$R \frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T)$$

insolation      albedo      re-radiation      transport

$Q = 343 \text{ W/m}^2 \quad A = 202 \text{ W/m}^2$   
 $B = 1.9 \text{ W/m}^2/\text{C} \quad C = 3.04 \text{ W/m}^2/\text{C}$   
 $\alpha_1 = 0.32$ : (water and land)  
 $\alpha_2 = 0.62$ : (ice)

$y$  = proportion of Earth's surface between latitude  $-\arcsin(y)$  and  $+\arcsin(y)$ ,  
 $\eta$  = proportion of Earth's surface that is ice-free.

If  $\sigma$  = surface area of Earth ( $\approx 5.1 \times 10^{14} \text{ m}^2$ ), then  
 $\sigma(1 - \eta)$  = surface area of ice in square meters.

## Ice-Albedo Feedback

**Budyko-Sellers Model**

$$R \frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T)$$

insolation      albedo      re-radiation      transport

### Heat Capacity

$R$  = heat capacity of Earth's surface ( $\text{J/m}^2/\text{C}$ ).  
Heat capacity of water  $\approx 4 \text{ J/g}^\circ\text{C}$ .  
Mass of  $1 \text{ cm}^3$  of  $\text{H}_2\text{O} \approx 1 \text{ g}$ .

Assumption: Earth's surface consists of  $100 \text{ m}$  of  $\text{H}_2\text{O}$ .  
Each square meter of Earth's surface consists of  $100 \text{ m}^3 = 10^8 \text{ cm}^3$  of  $\text{H}_2\text{O}$ , or about  $10^8 \text{ g}$ , so  
 $[R \approx 4 \times 10^8 \text{ J/m}^2/\text{C}]$

## Ice-Albedo Feedback

**Budyko-Sellers Model**

$$R \frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T)$$

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### Time Scale

A Watt is a Joule per second, but all quantities are annual averages.  
One year is approximately  $\kappa = 3.16 \times 10^7$  seconds.  
If we measure time in years instead of seconds, the equation becomes

$$\left[ \frac{\partial T}{\partial t} = \frac{\kappa}{R} (Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T)) \right]$$



## Ice-Albedo Feedback

### Budyko-Sellers Model

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Q_s(y)(1-\alpha(y,\eta)) - (A + BT) + C(\bar{T} - T))$$

**Dynamics of the Ice Line**

Widiasih equation:

$$\frac{d\eta}{dt} = \varepsilon(T_b - T_c)$$

where

$$T_b = \frac{T(\eta-) + T(\eta+)}{2}, \quad T_c = -10^\circ\text{C}$$

$$\frac{d\eta}{dt} = \varepsilon\kappa(T_b - T_c)$$

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Q_s(y)(1-\alpha(y,\eta)) - (A + BT) + C(\bar{T} - T))$$



## Ice-Albedo Feedback

### Budyko-Sellers-Widiasih Model

$$\frac{d\eta}{dt} = \varepsilon\kappa(T_b - T_c)$$

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Q_s(y)(1-\alpha(y,\eta)) - (A + BT) + C(\bar{T} - T))$$

$$T_b = \frac{T(\eta-) + T(\eta+)}{2}$$

$$\bar{T} = \bar{T}(t) = \int_0^t T(y,t) dy$$

**What about  $\varepsilon$  ?**

The parameter  $\varepsilon$  should be determined from climate data.  
(good project!)

Best current guess:  
 $\varepsilon \approx 10^{-12}$



## Ice-Albedo Feedback

### Budyko-Sellers-Widiasih Model

**Heat of Fusion**

$$\frac{d\eta}{dt} = \varepsilon\kappa(T_b - T_c)$$

To move the ice line, we must melt or freeze water.  
Heat of fusion of water: 334 J/g , or  $3.34 \times 10^8 \text{ J/m}^3$

Assumption: Average thickness of ice is 450 m.  
Energy to melt 1 m<sup>2</sup> of ice:  $\Omega = 1.5 \times 10^{11} \text{ Joules}$   
Energy to move ice line from  $\eta$  to  $\eta + \Delta\eta$  :

$$\Omega\Delta\eta \text{ Joules, or } \Omega\Delta\eta \text{ J/m}^2$$

( $\sigma$  = surface area of Earth.)

Energy needed to move the ice line:

$$\Omega \frac{d\eta}{dt} \text{ J/yr/m}^2$$



## Ice-Albedo Feedback

### Budyko-Sellers-Widiasih Model

$$\frac{d\eta}{dt} = \varepsilon\kappa(T_b - T_c)$$

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Q_s(y)(1-\alpha(y,\eta)) - (A + BT) + C(\bar{T} - T))$$

Energy to raise surface temperature:  $R \frac{\partial T}{\partial t} \text{ J/yr/m}^2$

Energy to move ice line:  $\Omega \frac{d\eta}{dt} \text{ J/yr/m}^2$

$$R \frac{\partial T}{\partial t} + \Omega \frac{d\eta}{dt} = \kappa(Q_s(y)(1-\alpha(y,\eta)) - (A + BT) + C(\bar{T} - T))$$

Altogether:

$$\frac{d\eta}{dt} = \varepsilon\kappa(T_b - T_c)$$

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Q_s(y)(1-\alpha(y,\eta)) - (A + BT) + C(\bar{T} - T) - \varepsilon\Omega(T_b - T_c))$$



## Ice-Albedo Feedback

### Budyko-Sellers-Widiasih Model

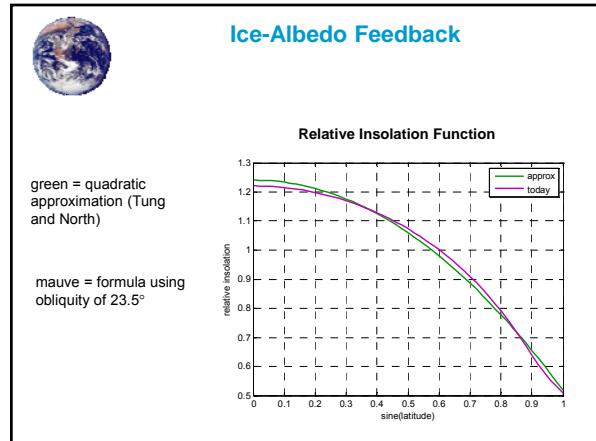
$$\frac{d\eta}{dt} = \varepsilon\kappa(T_b - T_c)$$

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Q_s(y)(1-\alpha(y,\eta)) - (A + BT) + C(\bar{T} - T) - \varepsilon\Omega(T_b - T_c))$$

**State Space**

$\eta$  lives in  $[0, 1]$ .  
 $T$  lives in a space of functions on  $[0, 1]$ .

**What space of functions?**



### Ice-Albedo Feedback

If  $s$  is assumed to be quadratic, then the equilibrium solutions are piecewise quadratic with a discontinuity at the ice boundary.

Two equilibrium solutions:  
 small cap: stable  
 large cap: unstable

Not equilibria:  
 ice free  
 snowball

equilibrium temperature profiles

temperature (°C)

sin(latitude)

— ice free — snowball — small cap — big cap

### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

$$\frac{d\eta}{dt} = \varepsilon K(T_b - T_c)$$

$$\frac{\partial U}{\partial t} = \frac{\kappa}{R} (Qs(y)(1-\alpha_i) - (A+BU) + C(\bar{T}-U) - \varepsilon\Omega(T_b - T_c)), \quad y < \eta$$

$$\frac{\partial V}{\partial t} = \frac{\kappa}{R} (Qs(y)(1-\alpha_2) - (A+BV) + C(\bar{T}-V) - \varepsilon\Omega(T_b - T_c)), \quad y > \eta$$

**State Space**

**What space of functions?**

Assume that  $T$  lives in the space of functions on  $[0, 1]$  which are piecewise quadratic with a discontinuity at the ice boundary.

$$T(y) = \begin{cases} U(y), & y < \eta, \\ V(y), & y > \eta, \\ (U(\eta) + V(\eta))/2, & y = \eta, \end{cases}$$

where  $U$  and  $V$  are quadratic on  $[0, 1]$ .

### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

$$\frac{d\eta}{dt} = \varepsilon K(T_b - T_c)$$

$$\frac{\partial U}{\partial t} = \frac{\kappa}{R} (Qs(y)(1-\alpha_i) - (A+BU) + C(\bar{T}-U) - \varepsilon\Omega(T_b - T_c)), \quad y < \eta$$

$$\frac{\partial V}{\partial t} = \frac{\kappa}{R} (Qs(y)(1-\alpha_2) - (A+BV) + C(\bar{T}-V) - \varepsilon\Omega(T_b - T_c)), \quad y > \eta$$

Ice boundary temperature:

$$T_b = \frac{U(\eta) + V(\eta)}{2}$$

Global mean temperature:

$$\bar{T} = \int_0^\eta U(y, t) dy + \int_\eta^1 V(y, t) dy$$

### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

**Legendre Polynomials**

Since  $s$  is even, we assume that  $U$  and  $V$  are even, and expand using the first two even Legendre polynomials:

$$p_0(y) = 1$$

$$p_2(y) = \frac{1}{2}(3y^2 - 1)$$

$$U(y, t) = u_0(t)p_0(y) + u_2(t)p_2(y) = u_0(t) + u_2(t)p_2(y),$$

$$V(y, t) = v_0(t)p_0(y) + v_2(t)p_2(y) = v_0(t) + v_2(t)p_2(y),$$

$$s(y) = s_0p_0(y) + s_2p_2(y) = 1 + s_2p_2(y).$$

$$\frac{\partial U}{\partial t} = \dot{u}_0 + \dot{u}_2 p_2(y), \quad \left( \dot{u} = \frac{du}{dt} \right)$$

$$\frac{\partial V}{\partial t} = \dot{v}_0 + \dot{v}_2 p_2(y).$$

### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

$$\frac{\partial U}{\partial t} = \frac{\kappa}{R} (Qs(y)(1-\alpha_i) - (A+BU) + C(\bar{T}-U) - \varepsilon\Omega(T_b - T_c))$$

$$\frac{\partial V}{\partial t} = \frac{\kappa}{R} (Qs(y)(1-\alpha_2) - (A+BV) + C(\bar{T}-V) - \varepsilon\Omega(T_b - T_c))$$

becomes

$$\dot{u}_0 + \dot{u}_2 p_2(y) = \frac{\kappa}{R} (Q(1+s_2p_2(y))(1-\alpha_i) - A + C\bar{T} - (B+C)(u_0 + u_2p_2(y)) - \varepsilon\Omega(T_b - T_c))$$

$$\dot{v}_0 + \dot{v}_2 p_2(y) = \frac{\kappa}{R} (Q(1+s_2p_2(y))(1-\alpha_2) - A + C\bar{T} - (B+C)(v_0 + v_2p_2(y)) - \varepsilon\Omega(T_b - T_c))$$

Equate coefficients:

$$\dot{u}_0 = \frac{\kappa}{R} (Q(1-\alpha_i) - A + C\bar{T} - (B+C)u_0 - \varepsilon\Omega(T_b - T_c))$$

$$\dot{u}_2 = \frac{\kappa}{R} (Qs_2(1-\alpha_i) - (B+C)u_2)$$

$$\dot{v}_0 = \frac{\kappa}{R} (Q(1-\alpha_2) - A + C\bar{T} - (B+C)v_0 - \varepsilon\Omega(T_b - T_c))$$

$$\dot{v}_2 = \frac{\kappa}{R} (Qs_2(1-\alpha_2) - (B+C)v_2)$$

### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

**Global Mean Temperature**

$$\bar{T} = \int_0^1 T(y) dy = \int_0^\eta U(y) dy + \int_\eta^1 V(y) dy$$

$$= \int_0^\eta (u_0 + u_2 p_2(y)) dy + \int_\eta^1 (v_0 + v_2 p_2(y)) dy$$

$$= \eta u_0 + (1-\eta) v_0 + u_2 \int_0^\eta p_2(y) dy + v_2 \int_\eta^1 p_2(y) dy$$

Recall:  $p_2(y) = \frac{1}{2}(3y^2 - 1)$

Hence:

$$\int_0^\eta p_2(y) dy = \frac{1}{2}(\eta^3 - \eta) = P_2(\eta), \quad \int_\eta^1 p_2(y) dy = -P_2(\eta)$$

$$\bar{T} = \eta u_0 + (1-\eta) v_0 + P_2(\eta)(u_2 - v_2)$$



### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

**Ice Line Temperature**

$$T_b = \frac{U(\eta) + V(\eta)}{2}$$

$$= \frac{1}{2} (u_0 + u_2 p_2(\eta) + v_0 + v_2 p_2(\eta))$$

$$T_b = \frac{u_0 + v_0}{2} + \frac{u_2 + v_2}{2} p_2(\eta)$$


### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

**Summary**

$$\dot{\eta} = \varepsilon K(T_b - T_c)$$

$$\dot{u}_0 = \frac{K}{R} (Q(1-\alpha_1) - A + C\bar{T} - (B+C)u_0 - \varepsilon\Omega(T_b - T_c))$$

$$\dot{v}_0 = \frac{K}{R} (Q(1-\alpha_2) - A + C\bar{T} - (B+C)v_0 - \varepsilon\Omega(T_b - T_c))$$

$$\dot{u}_2 = \frac{K}{R} (Qs_2(1-\alpha_1) - (B+C)u_2)$$

$$\dot{v}_2 = \frac{K}{R} (Qs_2(1-\alpha_2) - (B+C)v_2)$$

$$T_b = \frac{u_0 + v_0}{2} + \frac{u_2 + v_2}{2} p_2(\eta)$$

$$\bar{T} = \eta u_0 + (1-\eta)v_0 + P_2(\eta)(u_2 - v_2)$$


### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

**Quadratic Modes**

$$\dot{u}_2 = \frac{K}{R} (Qs_2(1-\alpha_1) - (B+C)u_2)$$

$$\dot{v}_2 = \frac{K}{R} (Qs_2(1-\alpha_2) - (B+C)v_2)$$

The dynamics of  $u_2$  and  $v_2$  are independent of each other and of the other variables (including  $\eta$ ). These variables decay exponentially to

$$u_2^* = \frac{Qs_2(1-\alpha_1)}{(B+C)} \quad v_2^* = \frac{Qs_2(1-\alpha_2)}{(B+C)}$$

Exercise: All higher order modes behave the same way.



### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

**Globally Attracting 3D Subspace**

$$\dot{\eta} = \varepsilon K(T_b - T_c)$$

$$\dot{u}_0 = \frac{K}{R} (Q(1-\alpha_1) - A + C\bar{T} - (B+C)u_0 - \varepsilon\Omega(T_b - T_c))$$

$$\dot{v}_0 = \frac{K}{R} (Q(1-\alpha_2) - A + C\bar{T} - (B+C)v_0 - \varepsilon\Omega(T_b - T_c))$$

where

$$T_b = \frac{u_0 + v_0 + \frac{u_2^* + v_2^*}{2} p_2(\eta)}{2} = \frac{u_0 + v_0}{2} + p_2(\eta) \frac{Qs_2(1-\alpha_0)}{B+C}, \quad \alpha_0 = \frac{\alpha_1 + \alpha_2}{2}$$

$$\bar{T} = \eta u_0 + (1-\eta)v_0 + P_2(\eta)(u_2^* - v_2^*)$$

$$= \eta u_0 + (1-\eta)v_0 + P_2(\eta) \frac{Qs_2(\alpha_2 - \alpha_1)}{B+C}$$


### Ice-Albedo Feedback

**Budyko-Sellers-Widiasih Model**

**Globally Attracting 3D Subspace**

**Summary**

$$\dot{\eta} = \varepsilon K(T_b - T_c)$$

$$\dot{u}_0 = \frac{K}{R} (Q(1-\alpha_1) - A + C\bar{T} - (B+C)u_0 - \varepsilon\Omega(T_b - T_c))$$

$$\dot{v}_0 = \frac{K}{R} (Q(1-\alpha_2) - A + C\bar{T} - (B+C)v_0 - \varepsilon\Omega(T_b - T_c))$$

**Ice Boundary Temperature**

$$T_b = \frac{u_0 + v_0}{2} + p_2(\eta) \frac{Qs_2(1-\alpha_0)}{B+C}$$

**Global Mean Temperature**

$$\bar{T} = \eta u_0 + (1-\eta)v_0 + P_2(\eta) \frac{Qs_2(\alpha_2 - \alpha_1)}{B+C}$$