

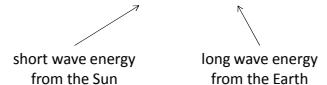
The Earth's Energy Imbalance
 Richard McGehee
 School of Mathematics
 University of Minnesota
 Mathematics of Climate Seminar
 October 6, 2015



Energy Imbalance

Conservation of Energy

temperature change \sim energy in – energy out



short wave energy from the Sun

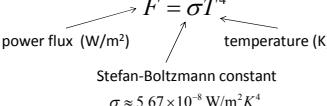
long wave energy from the Earth

Everything else is detail.

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Energy Imbalance

Stefan-Boltzmann Law



power flux (W/m²)

temperature (K)

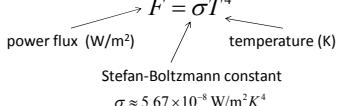
Stefan-Boltzmann constant

$\sigma \approx 5.67 \times 10^{-8} \text{ W/m}^2 \text{ K}^4$

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Energy Imbalance

Stefan-Boltzmann Law



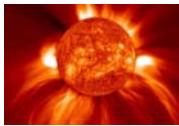
power flux (W/m²)

temperature (K)

Stefan-Boltzmann constant

$\sigma \approx 5.67 \times 10^{-8} \text{ W/m}^2 \text{ K}^4$

Example



surface temperature of the Sun: 5780K
 power flux: $5.67 \times 10^{-8} \times (5780)^4 = 6.33 \times 10^7 \text{ W/m}^2$
 total solar power output: $6.33 \times 10^7 \times 4\pi(r_s)^2$, where r_s = radius of the sun = $6.96 \times 10^8 \text{ m}$
 total solar output: $3.85 \times 10^{26} \text{ W}$
 200 nanoseconds = time it takes for the sun to produce the equivalent of the annual global electricity production ($7.3 \times 10^{19} \text{ Joules}$)

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Energy Imbalance

Insolation

Solar flux at a distance r from the sun:

$$F = \frac{6.33 \times 10^7 \cdot 4\pi r_s^2}{4\pi r^2} = 6.33 \times 10^7 \left(\frac{r_s}{r}\right)^2 \text{ W/m}^2$$

$$r_s = 6.96 \times 10^8 \text{ m}$$

$$r = 1.5 \times 10^{11} \text{ m}$$

$$F = 1368 \text{ W/m}^2$$

Power intercepted by the Earth:

$$F \times \pi r_E^2 \text{ Watts}, \quad r_E = \text{radius of Earth} = 6.37 \times 10^6 \text{ m}$$

$$F = 1.74 \times 10^{17} \text{ W}$$

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Energy Imbalance

Insolation

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Biologically Stored Energy

total coal reserves: 10^{15} kg
 energy content: $3 \times 10^7 \text{ J/kg}$
 total energy in coal reserves: $3 \times 10^{22} \text{ J}$
 $= 2 \text{ days of insolation}$



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Energy Imbalance



Insolation

Global Average Insolation
intercepted flux: $F = 1368 \text{ W/m}^2$
Earth cross-section: πr_E^2
surface area: $4\pi r_E^2$
average flux: $1368/4 = 342 \text{ W/m}^2 = Q$

Simple Model
Assume that Earth is a perfectly thermally conducting black body.

$$Q = \sigma T^4$$

$$T = (Q/\sigma)^{1/4} = (342/5.67 \times 10^{-8})^{1/4}$$

$$= 279 \text{ K} = 6^\circ \text{C} = 43^\circ \text{F}$$

Dynamics

$$R \frac{dT}{dt} = Q - \sigma T^4$$

heat capacity → stable equilibrium

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Energy Imbalance



Albedo

Not all the insolation reaches the surface. Some is reflected back into space.
The proportion reflected is called the albedo, denoted α .
For Earth, $\alpha \approx 0.3$.

Simple Model
Assume that Earth is a perfectly thermally conducting black body, but only 70% of the insolation is absorbed.

$$T = (0.7 \cdot F/\sigma)^{1/4} = (0.7 \cdot 342/5.67 \times 10^{-8})^{1/4}$$

$$= 255 \text{ K} = -18^\circ \text{C} = 0^\circ \text{F}$$

Dynamics

$$R \frac{dT}{dt} = Q(1-\alpha) - \sigma T^4$$

stable equilibrium

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Energy Imbalance



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Why isn't the Earth a snowball?



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Energy Imbalance

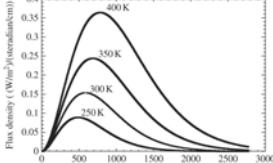


Black Body Radiation

Planck's Function

$$\text{flux density} \longrightarrow B(\nu, T) = \frac{2h\nu^3}{c^2} \frac{1}{e^{h\nu/kT} - 1}$$

frequency → Planck's constant
temperature → speed of light
Boltzmann's constant



wave number = $\frac{\nu}{c}$

Raymond T. Pierrehumbert, *Principles of Planetary Climate*, Cambridge University Press, 2010.

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Energy Imbalance



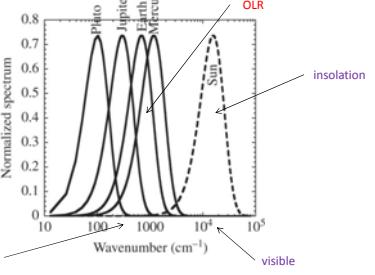
Insolation vs. OLR

OLR = Outgoing Longwave Radiation

Normalized spectrum

Wavenumber (cm^{-1})

infrared → visible



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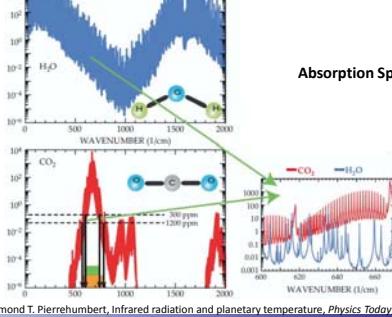
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Energy Imbalance



Greenhouse Gases

Absorption Spectra



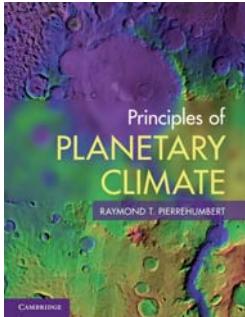
Raymond T. Pierrehumbert, *Infrared radiation and planetary temperature*, Physics Today, 2011.

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Energy Imbalance

Greenhouse Effect



How to learn more.

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Energy Imbalance

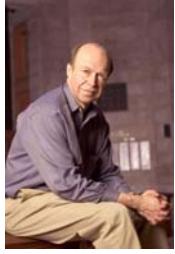
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Average Insolation: 342 Watts per square meter

Current Heat Imbalance:
 $0.85 \pm 0.15 \text{ W m}^{-2}$

Unit of energy: Watt-year (W yr)
1 W yr = $3.15 \times 10^7 \text{ J}$
 $= 8761 \text{ W hr} = 8.761 \text{ kWh}$

Yearly Insolation:
 $342 \text{ W yr m}^{-2} = 1.08 \times 10^{10} \text{ J}$



https://en.wikipedia.org/wiki/James_Hansen

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Ice sheet melting to raise sea level 1 meter. Assume ice starts at -10°C and ends at mean ocean surface temperature ($+15^\circ\text{C}$). Energy required is 100 cal/g (80 cal/g for melting). Energy for 1 meter of sea level: $100 \text{ g/cm}^2 \times 100 \text{ cal/g} \times 4.19 \text{ joules/cal} \times \text{area Earth} = 0.7 - 1.5 \times 10^{27} \text{ joules} = 9.3 \text{ W yr}^{-1}$

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Energy Imbalance

Assume all insolation goes toward warming.

Energy to warm air 1°C : 0.32 W yr m^{-2}
Power from Sun: 342 W m^{-2}
Time: $0.32/342 = 0.00094 \text{ yr} = 8.2 \text{ hr}$

Assume a global heat imbalance of 0.85 W m^{-2} .

Energy to warm air 1°C : 0.32 W yr m^{-2}
Heat imbalance: 0.85 W m^{-2}
Time: $0.32/0.85 = 0.376 \text{ yr} = 4.5 \text{ months}$

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Energy Imbalance

Assume all insolation goes toward warming.

Energy to warm land surface 1°C : 0.7 W yr m^{-2}
Power from Sun: 342 W m^{-2}
Time: $0.7/342 = 0.0020 \text{ yr} = 18 \text{ hr}$

Assume a global heat imbalance of 0.85 W m^{-2} .

Energy to warm land surface 1°C : 0.7 W yr m^{-2}
Heat imbalance: 0.85 W m^{-2}
Time: $0.7/0.85 = 0.376 \text{ yr} = 9.9 \text{ months}$

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Land surface warming by 1°C. The depth of penetration of a thermal wave into the Earth's crust is 10 years, weighted by AT, is $\sim 10 \text{ m}$. With density = 3 g/cm^3 , heat capacity = $0.2 \text{ cal/g}^\circ\text{C}$, and 0.29 fractional land coverage, land heat storage is $10^9 \text{ cm}^3 \times 3 \text{ g/cm}^3 \times 0.2 \text{ cal/g}^\circ\text{C} \times 1^\circ\text{C} \times 4.19 \text{ joules/cal} \times \text{area Earth} = 0.29 - 0.37 \times 10^{21} \text{ joules} = 0.23 \text{ W yr}^{-1}$. [In a century the depth of penetration is ~ 3 times more than in a decade, so heat storage in a century due to 1°C warming is $\sim 0.7 \text{ W yr}^{-1}$]

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 **Energy Imbalance**

Assume all insolation goes toward warming.

Energy to melt all the sea ice: 1.3 W yr m^{-2}
Power from Sun: 342 W m^{-2}
Time: $1.3/342 = 0.0038 \text{ yr} = 33 \text{ hr}$

Assume a global heat imbalance of 0.85 W m^{-2} .

Energy to melt all the sea ice: 1.3 W yr m^{-2}
Heat imbalance: 0.85 W m^{-2}
Time: $1.3/0.85 = 1.53 \text{ yr} = 18 \text{ months}$

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Assume all insolation goes toward warming.

Energy to raise sea level 1 meter by melting ice sheets: 9.3 W yr m^{-2}
Power from Sun: 342 W m^{-2}
Time: $9.3/342 = 0.027 \text{ yr} = 9.9 \text{ days}$

Assume a global heat imbalance of 0.85 W m^{-2} .

Energy to raise sea level 1 meter by melting ice sheets: 9.3 W yr m^{-2}
Heat imbalance: 0.85 W m^{-2}
Time: $9.3/0.85 = 10.9 \text{ yr}$

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Assume all insolation goes toward warming.

Energy to warm ocean 1°C to depth 1 km: 93 W yr m^{-2}
Power from Sun: 342 W m^{-2}
Time: $93/342 = 0.27 \text{ yr} = 3.3 \text{ months}$

Assume a global heat imbalance of 0.85 W m^{-2} .

Energy to warm ocean 1°C to depth 1 km: 93 W yr m^{-2}
Heat imbalance: 0.85 W m^{-2}
Time: $93/0.85 = 109 \text{ years}$

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Energy Imbalance



What if we melt all the ice sheets?

All ice sheets will raise the sea level 70 m.
Will take $9.3 \text{ W yr m}^{-2} \times 70 = 650 \text{ W yr m}^{-2}$

Assume all insolation goes toward warming.

Energy to melting all ice sheets: 650 W yr m^{-2}
Power from Sun: 342 Wm^{-2}
Time: $650/342 = 1.9 \text{ years}$

Assume a global heat imbalance of 0.85 W m^{-2} .

Energy to melting all ice sheets: 650 W yr m^{-2}
Heat imbalance: 0.85 Wm^{-2}
Time: $650/0.85 = 765 \text{ years}$

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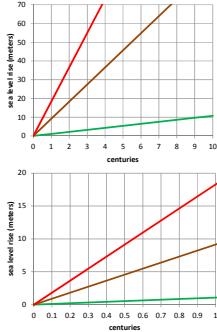
Energy Imbalance



Suppose that all the heat imbalance went to melting the glaciers.

It takes 9.3 Wyr/m^2 to turn glaciers into 1 meter of ocean. If the heat imbalance is $w \text{ W/m}^2$, the sea level would rise at the rate of $w/9.3 \text{ meters per year}$. At the current imbalance of 0.85 W/m^2 , the rate is about $0.109 \text{ meters per year}$, or $10.9 \text{ meters per century}$.

Melting all the glaciers would cause a sea level rise of about 70 meters and would take about 765 years at the current imbalance.



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Suppose now that all the heat imbalance first goes to raising the top kilometer of ocean by $0.5 \text{ }^\circ\text{C}$, and then goes to melting the glaciers.

It takes 46.5 Wyr/m^2 to raise the temperature of a kilometer of ocean by $0.5 \text{ }^\circ\text{C}$. If the heat imbalance is $w \text{ W/m}^2$, the increase would be achieved in $46.5/w$ years, after which the sea level would rise at $w/9.3 \text{ meters per year}$.

At the current imbalance of 0.85 W/m^2 , the ocean temperature increase would delay the sea level rise by about 55 years.

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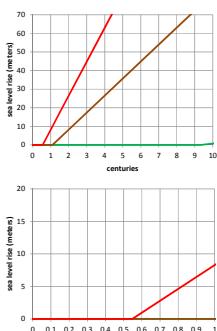
Energy Imbalance



Suppose instead that all the heat imbalance first goes to raising the top kilometer of ocean by $1 \text{ }^\circ\text{C}$, and then goes to melting the glaciers.

It takes 93 Wyr/m^2 to raise the temperature of a kilometer of ocean by $1 \text{ }^\circ\text{C}$. If the heat imbalance is $w \text{ W/m}^2$, the increase would be achieved in $93/w$ years, after which the sea level would rise at $w/9.3 \text{ meters per year}$.

At the current imbalance of 0.85 W/m^2 , the ocean temperature increase would delay the sea level rise by about 109 years.



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Energy Imbalance



Summary

Currently, it appears that the heat imbalance is mostly going to heating the ocean, not to melting ice. If this pattern continues, the danger for this century is more likely to come from weather changes than from sea level rise.

The current heat imbalance has the potential to raise the sea level by almost a meter per decade, a major threat to coastal cities worldwide.

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Energy Imbalance



Questions about the Coming Centuries

- How will the heat imbalance be divided between heating the ocean and melting the glaciers?
- How will the heat imbalance be affected by increasing atmospheric greenhouse gases?
- How will the heat imbalance be affected by increasing ocean temperatures?
- What happens to the weather as the ocean temperature rises and the ice caps melt?
- What should we do about coastal cities?

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Energy Imbalance

What can mathematicians do?

Models
Minimal complexity, aka "conceptual", "simple", "toy"
Intermediate complexity
Maximal complexity, aka "GCM"

Data
Parameter estimation (statistics)
Data assimilation

Quantification
Uncertainty
Resilience

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