

Math 5490
Topics in Applied Mathematics
Introduction to the Mathematics of Climate

Fall 2023
1:25 - 3:20 Tuesdays and Thursdays
Amundson Hall 162

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course website
 www-users.cse.umn.edu/~mcgehee/teaching/Math5490/

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Math 5490
Review

What determines the Earth's surface temperature?

Conservation of Energy
 Heat is a form of energy.
 Temperature measures heat.

temperature change ~ energy in - energy out

↑

short wave energy
from the Sun

↑

long wave energy
from the Earth

Everything else is detail.

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Black-Body Radiation
Stefan-Boltzmann Law

watts per square meter
power flux (W/m²)

$F = \sigma T^4$

kelvin
temperature (K)

0 K = -273°C = "absolute zero"

Stefan-Boltzmann constant
 $\sigma \approx 5.67 \times 10^{-8} \text{ W/m}^2\text{K}^4$


Reasonable approximation:
 Every body in the solar system radiates energy according to this law.

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What determines the Earth's surface temperature?

temperature change ~ energy in - energy out



short wave energy
from the Sun



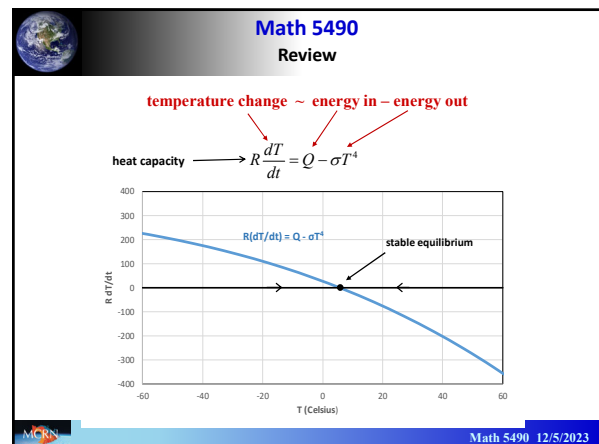
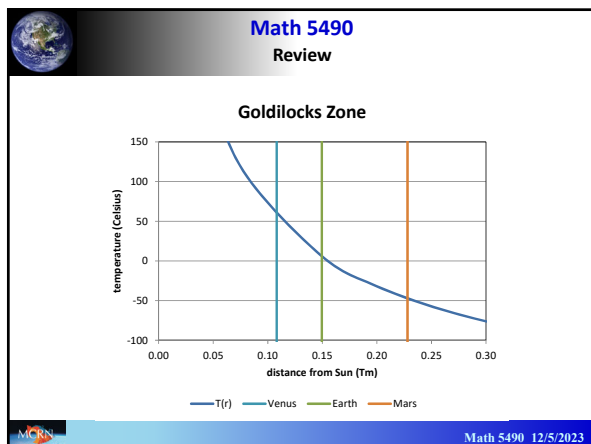
long wave energy
from the Earth

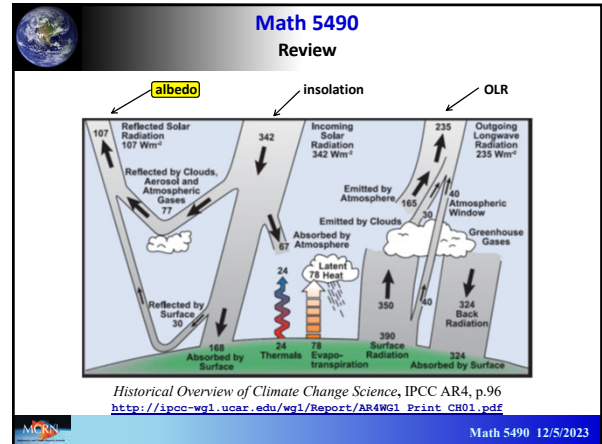
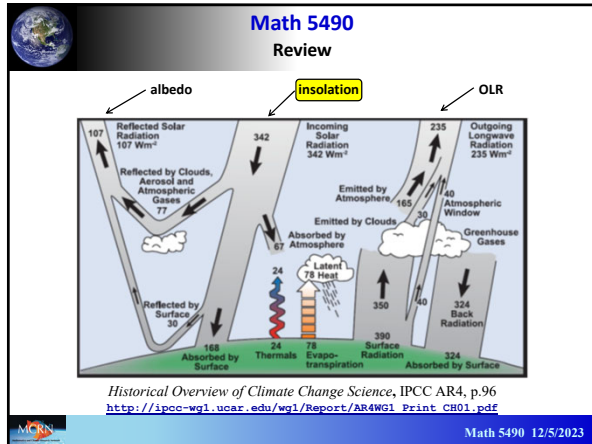
energy in from the Sun

At equilibrium, these are equal.

energy out from the Earth

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Albedo

Not all the insolation reaches the surface. Some is reflected back into space. The proportion reflected is called the albedo, denoted α . For Earth, $\alpha \approx 0.3$.

Simple Model
 Assume that Earth is a perfectly thermally conducting black body, but only 70% of the insolation is absorbed.

$$T = (0.7 \cdot Q / \sigma)^{1/4} = (0.7 \cdot 342 / 5.67 \times 10^{-8})^{1/4}$$

$$= 255\text{K} = -18^\circ\text{C} = 0^\circ\text{F}$$

Dynamics

$$R \frac{dT}{dt} = Q(1 - \alpha) - \sigma T^4$$

stable equilibrium

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Photosphere

A photosphere is the deepest region of a luminous object, usually a star, that is transparent to photons of certain wavelengths.

<https://en.wikipedia.org/wiki/Photosphere>

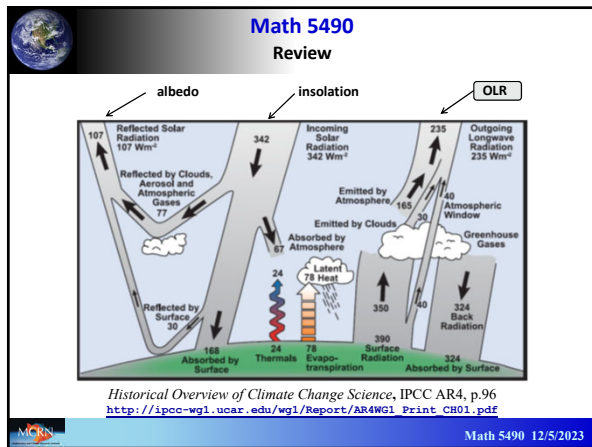
For the Earth, the photosphere is where the long wave photons escape into space. It is high in the atmosphere where the temperature is 255 K.

$$R \frac{dT}{dt} = Q(1 - \alpha) - \sigma T^4$$

T = photosphere temperature.

What about the surface temperature?

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OLR as a Function of Surface Temperature
 (Outgoing Longwave Radiation)

$$OLR \approx A + BT$$

A and B are determined from satellite observations.
 T is surface temperature (in Celsius).

$$A = 202 \text{ W/m}^2$$

$$B = 1.90 \text{ W/m}^2\text{K}$$

Dynamics

Kelvin $R \frac{dT}{dt} = Q(1 - \alpha) - \sigma T^4$ photosphere temperature

Celsius becomes $R \frac{dT}{dt} = Q(1 - \alpha) - (A + BT)$ global mean surface temperature

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OLR as a Function of Surface Temperature

$$OLR \approx A + BT$$

Important:
 $A + BT$ is not a linear approximation to the Stefan-Boltzmann equation.

Kelvin

$$R \frac{dT}{dt} = Q(1 - \alpha) - \sigma T^4$$

Dynamics

different

photosphere temperature

becomes

Celsius

$$R \frac{dT}{dt} = Q(1 - \alpha) - (A + BT)$$

becomes

global mean surface temperature

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Homogeneous Earth

$$R \frac{dT}{dt} = Q(1 - \alpha) - (A + BT)$$

What's missing?

Earth is not homogeneous. For example, it is warmer at the equator and colder at the poles. The temperature should depend on latitude.

Make T depend on $y = \sin(\text{latitude})$

$$R \frac{\partial T(y,t)}{\partial t} = Qs(y)(1 - \alpha) - (A + BT(y,t))$$

insolation distribution

$s(y) = \text{distribution across latitudes } \left(\int_0^1 s(y) dy = 1 \right)$

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Latitude Dependence

$$R \frac{\partial T(y,t)}{\partial t} = Qs(y)(1 - \alpha) - (A + BT(y,t))$$

$s(y) = \text{distribution across latitudes } \left(\int_0^1 s(y) dy = 1 \right)$

One can show that $\beta = \text{obliquity} = 23.4^\circ$

$$s(y) = \frac{2}{\pi^2} \int_0^{2\pi} \sqrt{1 - (y^2 \sin \beta \cos \theta - y \cos \beta)^2} d\theta$$

McGehee & Lehman, *SIAM J. Applied Dynamical Systems* **11** (2) (2012), 684–707.

Chylek and Coakley's quadratic approximation:
 $s(y) \approx 1 - 0.241(3y^2 - 1)$

Chylek & Coakley, *J. Atmos. Sci.* **32** (1975), 675–679.

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Insolation Distribution

green = quadratic approximation (Chylek & Coakley)

fuchsia = formula using obliquity of 23.5°

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What's Missing?

Thermohaline Circulation

Weather!

The second law of thermodynamics

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Budyko's Equation

$$R \frac{\partial T}{\partial t}(y,t) = Qs(y)(1 - \alpha) - (A + BT(y,t)) - C(\bar{T}(t) - T(y,t))$$

2nd law

global mean temperature $\bar{T}(t) = \int_0^1 T(y,t) dy$

$(\bar{T}(t) - T(y,t))$ interpretation

Each point on Earth's surface is trying to assume the global mean temperature. If the temperature at a point is below the global mean, then it heats up. If the temperature at that point is above the mean, then it cools off.

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Budyko's Equation


$$R \frac{\partial T}{\partial t} = Qs(y)(1-\alpha) - (A+BT) + C(\bar{T}-T)$$

There is still something missing. What role is played by the ice?

Ice-albedo Feedback

- temperature warms
- ice melts
- albedo decreases
- more sunlight absorbed
- temperature warms
- REPEAT

Why would it stop?



M. I. Budyko, "The effect of solar radiation variations on the climate of the Earth," *Tellus XXI*, 611-619, 1969.

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Budyko's Equation

$$R \frac{\partial T}{\partial t} = Qs(y)(1-\alpha) - (A+BT) + C(\bar{T}-T)$$

There is still something missing. Where is the ice?

Ice-albedo Feedback

- albedo of ice: 0.62
- albedo of land and water: 0.32

Assumption: there is a single boundary between ice and no ice occurring at $y = \eta$.

$$\alpha = \alpha(y, \eta) = \begin{cases} \alpha_1 = 0.32, & y < \eta \\ \alpha_2 = 0.62, & y > \eta \end{cases}$$

$$R \frac{\partial T}{\partial t} = Qs(y)(1-\alpha(y, \eta)) - (A+BT) + C(\bar{T}-T)$$

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Logical Fallacy

Climate changes occurred when humans didn't exist implies that humans are not the cause of climate change.

Reasonable Conclusion

Climate changes occurred when humans didn't exist suggests that we should try to understand why the climate changed in the past.

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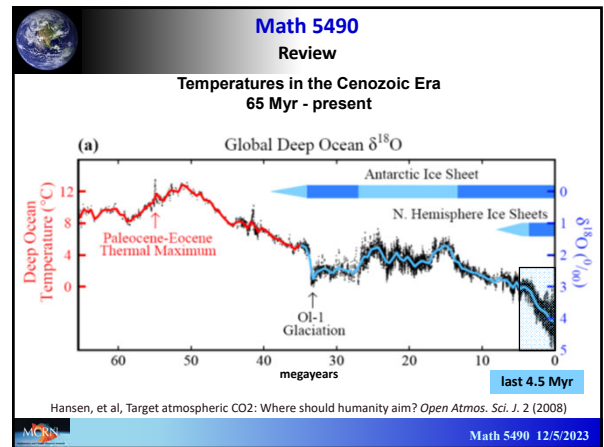
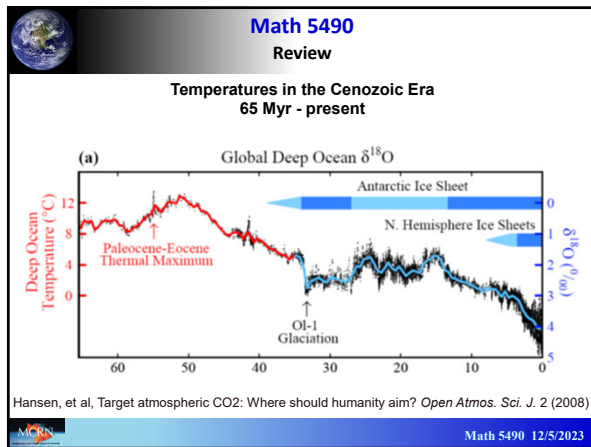
Earth's climate has changed many times in the past.

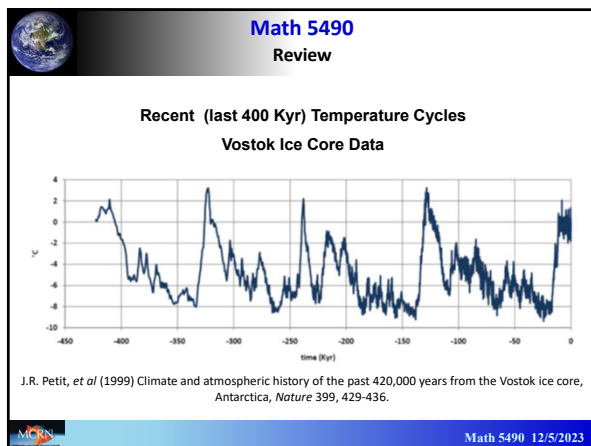
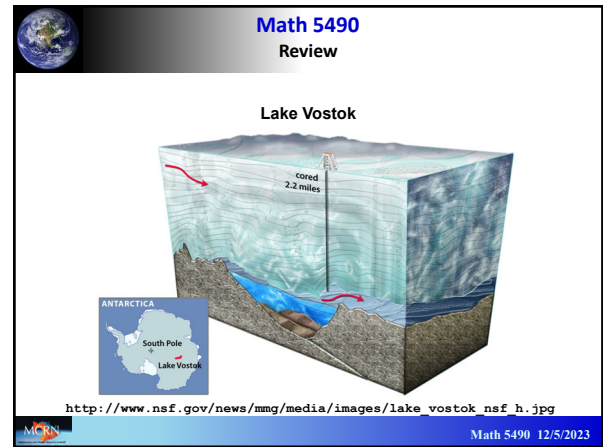
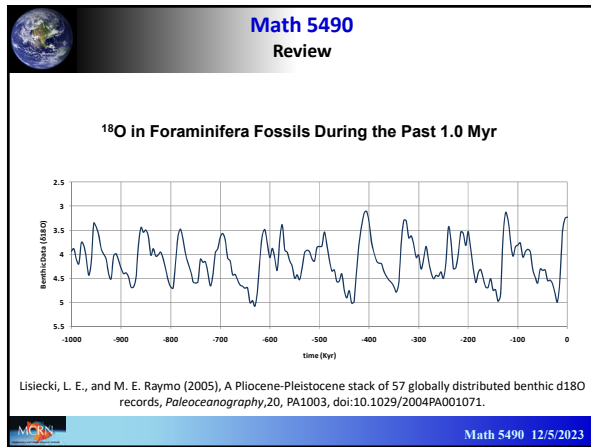
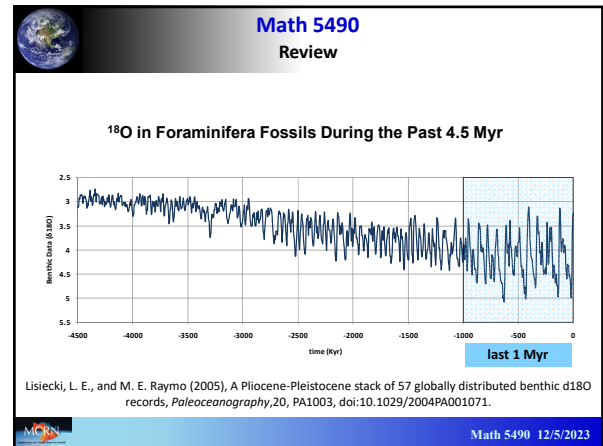
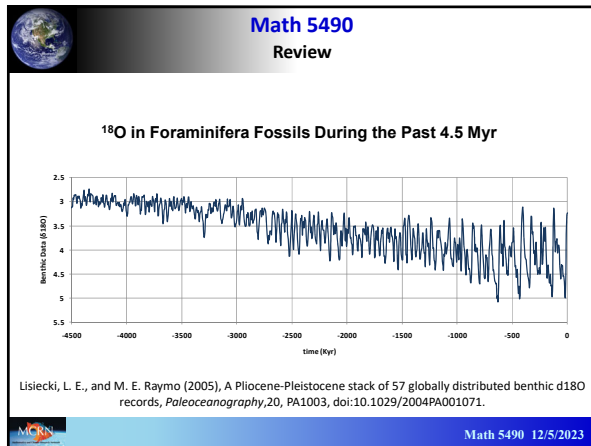
Paleoclimatology

Paleoclimatology is the scientific study of climates predating the invention of meteorological instruments, when no direct measurement data were available. As instrumental records only span a tiny part of Earth's history, the reconstruction of ancient climate is important to understand natural variation and the evolution of the current climate.

<https://en.wikipedia.org/wiki/Paleoclimatology>

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What Causes Glacial Cycles?

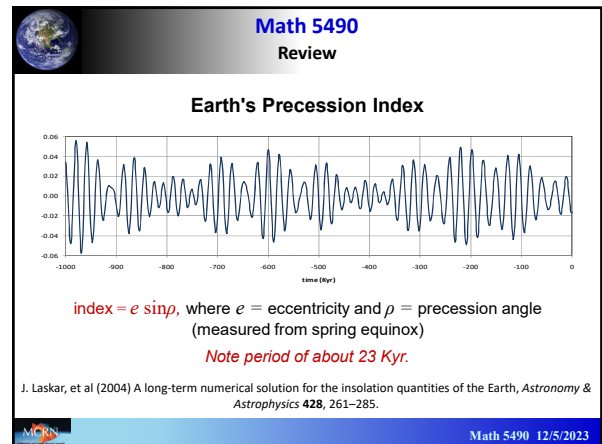
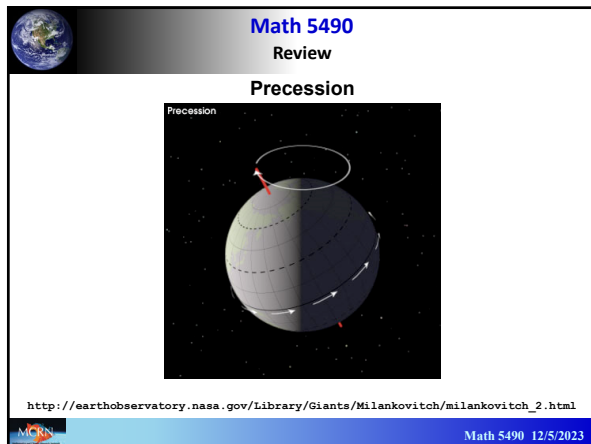
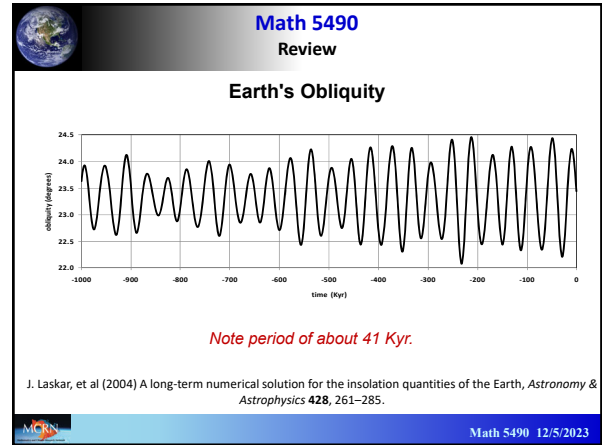
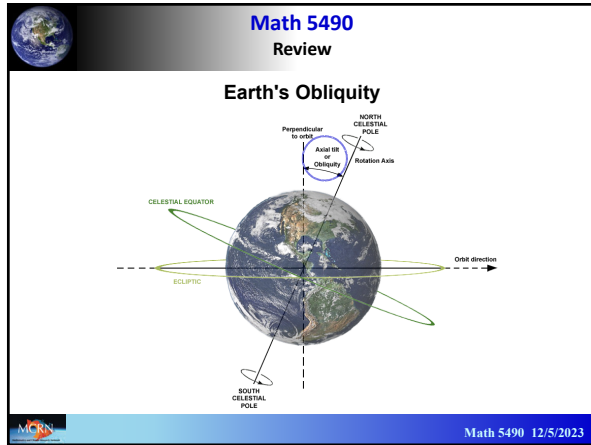
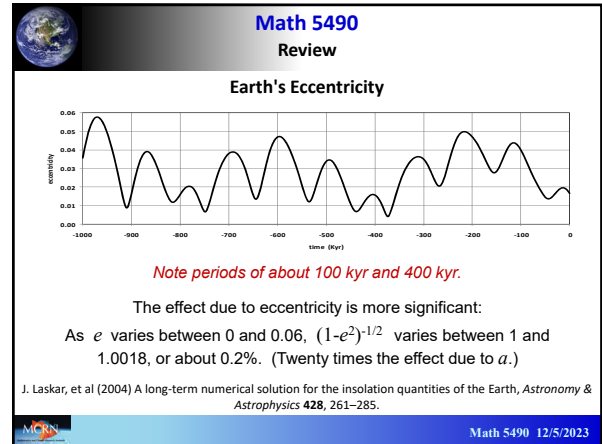
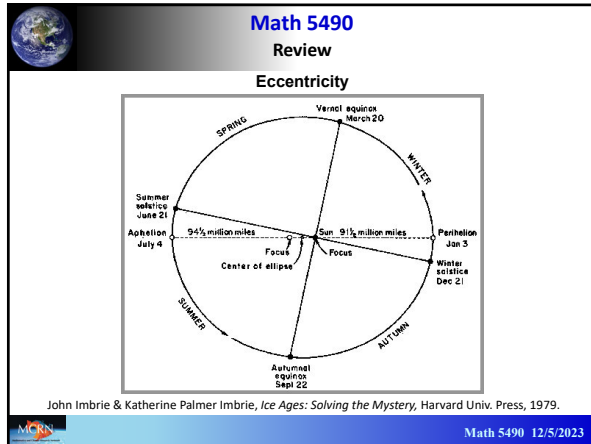
Widely Accepted Hypothesis

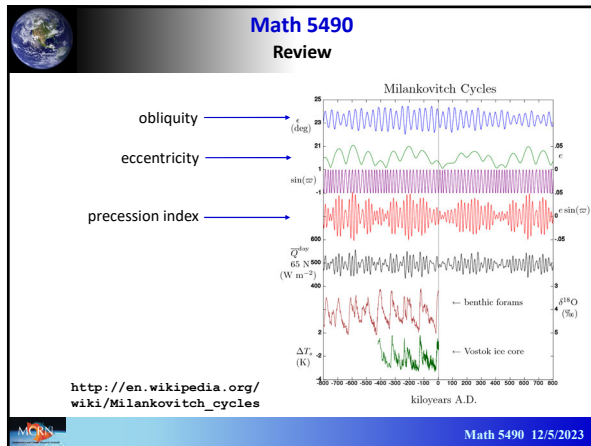
The glacial cycles are driven by the variations in the Earth's orbit (Milankovitch Cycles), causing a variation in incoming solar radiation (insolation).

This hypothesis is widely accepted, but also widely regarded as insufficient to explain the observations.

The additional hypothesis is that there are feedback mechanisms and/or triggering mechanisms that amplify the Milankovitch cycles. What these feedbacks are and how they work are not fully understood.

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Who was Milankovitch?

Milutin Milankovitch was a Serbian mathematician and professor at the University of Belgrade.

In 1920 he published his seminal work on the relation between insolation and the Earth's orbital parameters.

In 1941 he published a book explaining his entire theory.

His work was not fully accepted until 1976.

Milutin Milankovitch
1879-1958

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What happened in 1976?

Hays, Imbrie, and Shackleton, "Variations in the Earth's Orbit: Pacemaker of the Ice Ages," *Science* 194, 10 December 1976.

John Imbrie

James D. Hays

Nicholas Shackleton

"It is concluded that changes in the earth's orbital geometry are the fundamental cause of the succession of Quaternary ice ages."

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Hays, et al, Summary

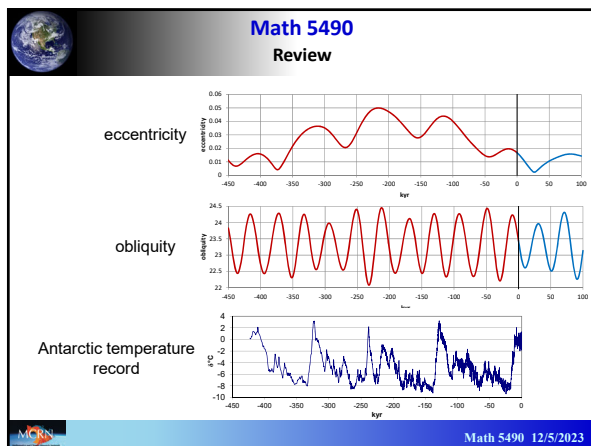
6) It is concluded that changes in the earth's orbital geometry are the fundamental cause of the succession of Quaternary ice ages.

7) A model of future climate based on the observed orbital-climate relationships, but ignoring anthropogenic effects, predicts that the long-term trend over the next seven thousand years is toward extensive Northern Hemisphere glaciation.

*Quoted by George Will, Washington Post, February 5, 2009

Hays, et al, *Science* 194 (1976), p. 1125

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Earth's climate has changed many times in the past.
Why do we think humans are responsible now?

Here's a reason.

The paleoclimate evidence points to the conclusion that the Earth should be entering a new ice age.

But we're not.

Instead, the Earth is warming and the ice sheets are melting.

The climate is not following the patterns of the last million years.

Something has changed.

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Earth's Carbon Cycle

Earth's climate has changed many times in the past.
 Why do we think humans are responsible now?
 Why do we think that atmospheric CO₂ has anything to do with climate change?
 Why do we think that the increase in atmospheric CO₂ has anything to do with human activity?
 Why do we think that atmospheric CO₂ has anything to do with surface temperature?

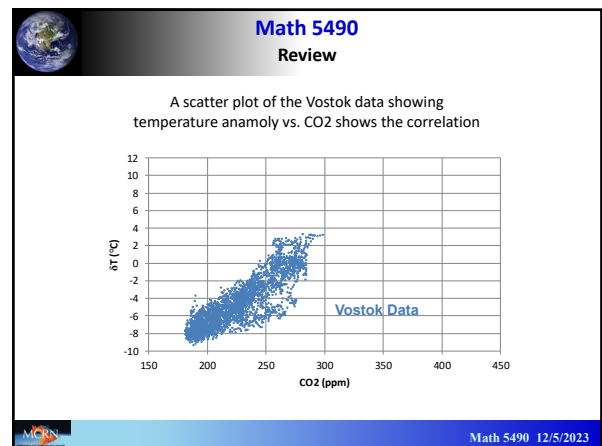
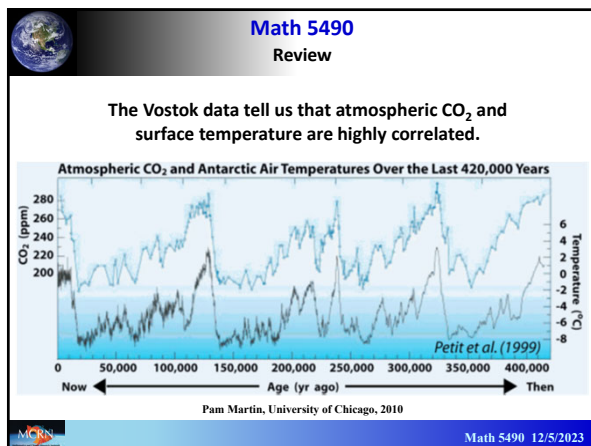
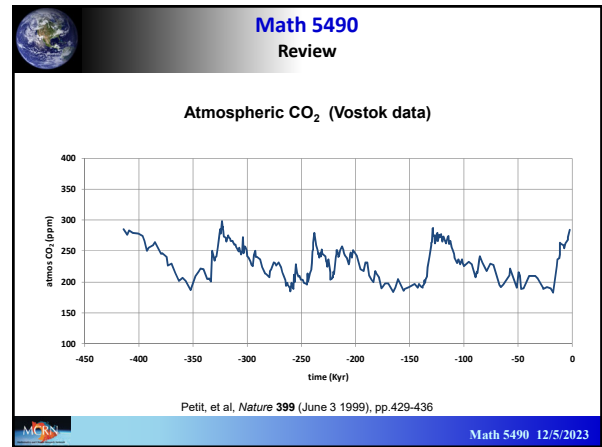
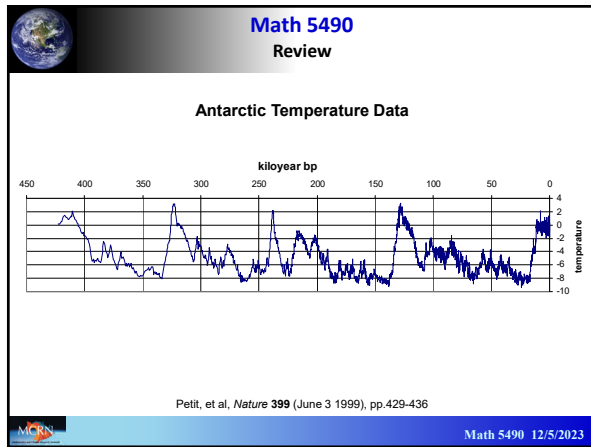
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Lake Vostok

http://www.nsf.gov/news/mmg/media/images/lake_vostok_nsf_h.jpg

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The Vostok data do not extend to the post-industrial times. Do we have anything recent?

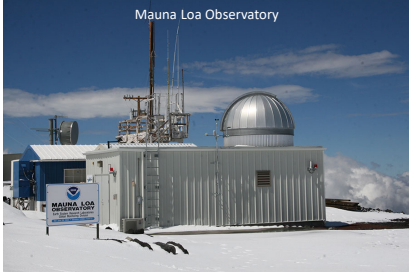
Yes!
The Keeling curve!

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The Keeling Curve

Mauna Loa Observatory



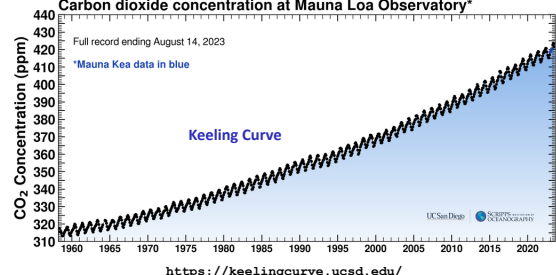
https://research.noaa.gov/Portals/0/EasyDNNnews/1502/200600p587EDNmain10061ml0_sign_miller.jpg

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Can we measure greenhouse gasses?

Carbon dioxide concentration at Mauna Loa Observatory*



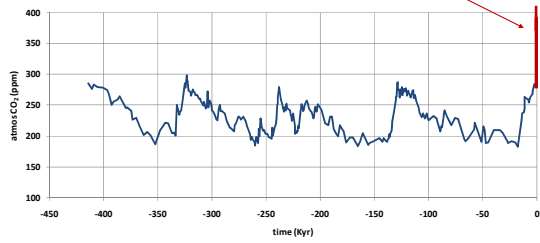
Full record ending August 14, 2023
*Mauna Kea data in blue
Keeling Curve
UC San Diego Scripps Institution of Oceanography
<https://keelingcurve.ucsd.edu/>

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Keeling curve on Vostok time scale.

Atmospheric CO₂ (Vostok data)



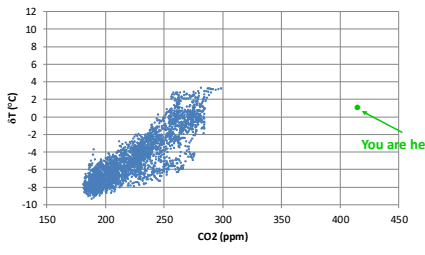
Petit, et al, *Nature* 399 (June 3 1999), pp.429-436

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Vostok Data

Current conditions are well outside the range recorded in the ice core data.



You are here.

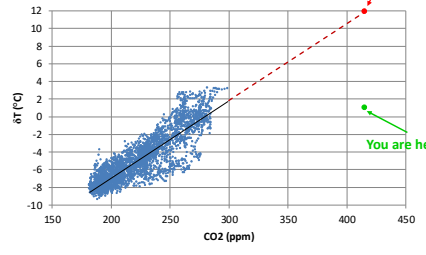
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Vostok Data

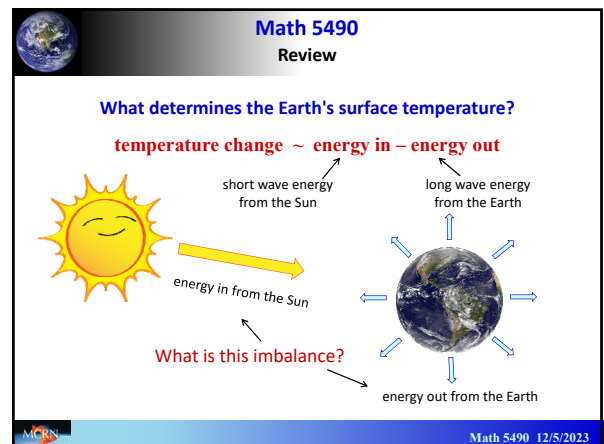
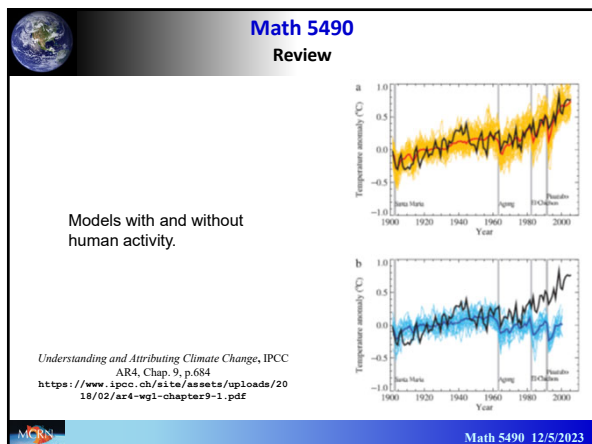
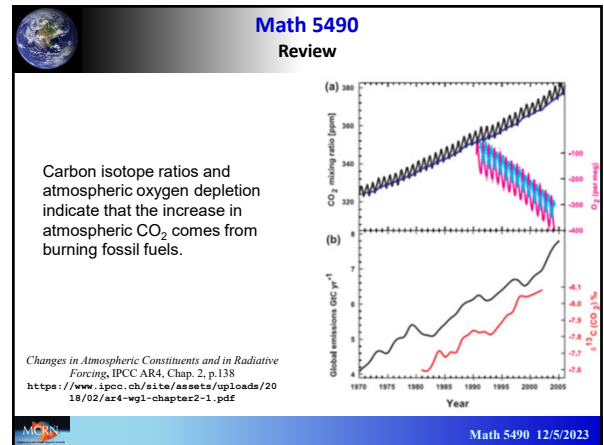
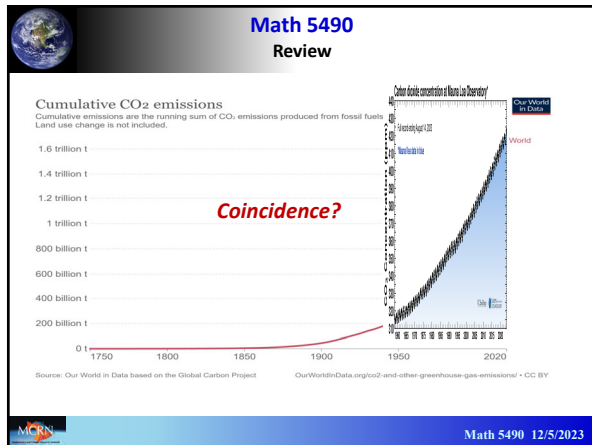
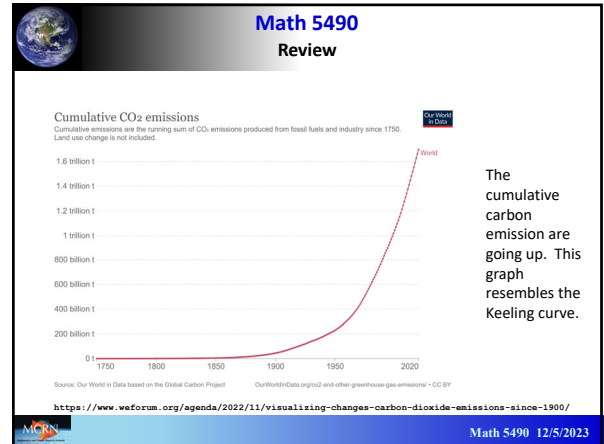
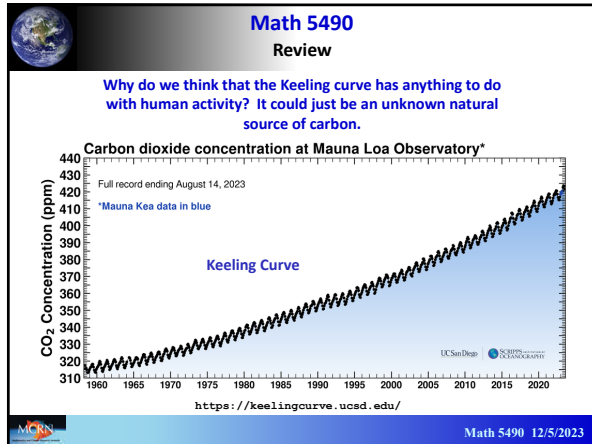
Extrapolate linear regression to 420 ppm CO₂.

Will you be here?



You are here.

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Earth's Heat Imbalance

Net TOA Radiation or Planetary Heat Uptake (Wm^{-2})

Year

— Net TOA Radiation (CERES)
— Planetary Heat Uptake (In Situ)

<https://www.nasa.gov/centers-and-facilities/langley/joint-nasa-noaa-study-finds-earths-energy-imbalance-has-doubled/>

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James Hansen

James Hansen arrested at a demonstration outside the White House, August 29, 2011

Hansen giving testimony before the United States Congress in 1988.

https://en.wikipedia.org/wiki/James_Hansen

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www.sciencemag.org SCIENCE VOL 308 3 JUNE 2005

RESEARCH ARTICLES

Earth's Energy Imbalance: Confirmation and Implications

James Hansen,^{1,2*} Larissa Nazarenko,^{3,2} Reto Ruedy,² Makiko Sato,^{3,2} Josh Willis,⁴ Anthony Del Genio,^{5,6} Dorothy Koch,^{1,2} Andrew Lacis,^{1,3} Ken Lo,⁷ Surabi Menon,⁸ Tica Novakov,⁹ Judith Perwitz,^{1,2} Gary Russell,¹ Gavin A. Schmidt,^{1,2} Nicholas Tausnev⁹

Our climate model, driven mainly by increasing human-made greenhouse gases and aerosols, among other forcings, calculates that Earth is now absorbing 0.85 ± 0.15 watts per square meter more energy from the Sun than it is emitting to space. This imbalance is confirmed by precise measurements of increasing ocean heat content over the past 30 years. Implications include (i) the expectation of additional global warming of about 0.6°C without further change of atmospheric composition; (ii) the confirmation of the climate system's lag in responding to forcings, implying the need for anticipatory actions to avoid any specified level of climate change; and (iii) the likelihood of acceleration of ice sheet disintegration and sea level rise.

equal forcing by CO₂ (9). F_p is an energy flux change arising in response to an imposed forcing agent. It is constant throughout the atmosphere, because it is evaluated after atmospheric temperature has been allowed to adjust to the presence of the forcing agent.

The largest forcing is due to well-mixed greenhouse gases (GHGs)—CO₂, CH₄, N₂O, CFCs (chlorofluorocarbons)—and other trace gases, adding 2.75 W/m² in 2003 relative to the 1880 value (Table 1). Ozone (O₃) and stratospheric H₂O from oxidation of increasing CH₄ bring the total GHG forcing to 3.00 W/m² (9). Estimated uncertainty in the total GHG forcing is ~15% (11, 12).

Atmospheric aerosols cause climate forcings by reflecting and absorbing radiation, as well as through indirect effects on cloud cover and cloud albedo (17). The aerosol scenario in our model uses estimated anthropogenic emissions from fuel use statistics and includes temporal changes in fossil-fuel use technologies (15). Our parameterization of aerosol

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Table S1. Planetary Heat Storage: Ocean, Ice, Air and Land.

Energy required to melt ice and warm the air, land and ocean by specified amounts.¹

Ocean warming by 1°C through 1 km depth of ocean. Heat storage is $1^\circ C \times 10^9 g/cm^3 \times 1 cal/g \times 4.19 joules/cal \times area Earth \times 0.7 - 15 \times 10^{22} joules - 93 W yr/m^2$.

Ice sheet melting to raise sea level 1 meter. Assume ice starts at -10°C and ends at mean ocean surface temperature (+15°C). Energy required is 100 cal/g (80 cal/g for melting). Energy for 1 meter of sea level: $100g/cm^3 \times 100cal/g \times 4.19 joules/cal \times area Earth \times 0.7 - 1.5 \times 10^{22} joules - 9.3 W yr/m^2$.

Sea ice melting (all sea ice on planet). Assume ice starts at -10°C and ends at mean ocean surface temperature (+15°C), and that sea ice covers 4% of the planet with mean thickness 2.5 m. Energy required is $250 g/cm^3 \times 100 cal/g (80 cal/g for melting) \times 4.19 joules/cal \times 0.04 \times area Earth - 2.14 \times 10^{22} joules - 1.3 W yr/m^2$.

Air warming by 1°C. The Earth's atmospheric mass is ~ 10 m of water. Heat capacity of air ~ 0.24 cal/g°C. Energy to raise air temperature 1°C: $1^\circ C \times 1000 g/cm^3 \times 0.24 cal/g^\circ C \times 4.19 joules/cal \times area Earth - 0.26 \times 10^{22} joules - 0.32 W yr/m^2$.

Land surface warming by 1°C. The depth of penetration of a thermal wave into the Earth's crust in 10 years, weighted by ΔT , is ~10 m. With density ~ 3 g/cm³, heat capacity ~ 0.2 cal/g°C, and 0.29 fractional land coverage, land heat storage is $10^3 cm^3/g/cm^3 \times 0.2 cal/g^\circ C \times 1^\circ C \times 4.19 joules/cal \times area Earth - 0.29 - 0.37 \times 10^{22} joules - 0.23 W yr$. [In a century the depth of penetration is ~3 times more than in a decade, so heat storage in a century due to 1°C warming is ~ 0.7 W yr/m².]

¹Note that 1 W sec = 1 joule, # sec/year = $\pi \times 10^7$, area Earth = $5.1 \times 10^{18} cm^2$, 1 W yr over full Earth ~ 1.61×10^{22} joules, ocean fraction of Earth ~ 0.7, 1 calorie = 4.19 joules.

James Hansen, et al, *Earth's Energy Imbalance: Confirmation and Implications*, SCIENCE 308 (2005), p. 1431

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Table S1. Planetary Heat Storage: Ocean, Ice, Air and Land.

Energy required to melt ice and warm the air, land and ocean by specified amounts.¹

Ocean warming by 1°C through 1 km depth of ocean. Heat storage is $1^\circ C \times 10^9 g/cm^3 \times 1 cal/g \times 4.19 joules/cal \times area Earth - 0.7 - 15 \times 10^{22} joules - 93 W yr/m^2$.

Ice sheet melting to raise sea level 1 meter. Assume ice starts at -10°C and ends at mean ocean surface temperature (+15°C). Energy required is 100 cal/g (80 cal/g for melting). Energy for 1 meter of sea level: $100g/cm^3 \times 100cal/g \times 4.19 joules/cal \times area Earth \times 0.7 - 1.5 \times 10^{22} joules - 9.3 W yr/m^2$.

Sea ice melting (all sea ice on planet). Assume ice starts at -10°C and ends at mean ocean surface temperature (+15°C), and that sea ice covers 4% of the planet with mean thickness 2.5 m. Energy required is $250 g/cm^3 \times 100 cal/g (80 cal/g for melting) \times 4.19 joules/cal \times 0.04 \times area Earth - 2.14 \times 10^{22} joules - 1.3 W yr/m^2$.

Air warming by 1°C. The Earth's atmospheric mass is ~ 10 m of water. Heat capacity of air ~ 0.24 cal/g°C. Energy to raise air temperature 1°C: $1^\circ C \times 1000 g/cm^3 \times 0.24 cal/g^\circ C \times 4.19 joules/cal \times area Earth - 0.26 \times 10^{22} joules - 0.32 W yr/m^2$.

Land surface warming by 1°C. The depth of penetration of a thermal wave into the Earth's crust in 10 years, weighted by ΔT , is ~10 m. With density ~ 3 g/cm³, heat capacity ~ 0.2 cal/g°C, and 0.29 fractional land coverage, land heat storage is $10^3 cm^3/g/cm^3 \times 0.2 cal/g^\circ C \times 1^\circ C \times 4.19 joules/cal \times area Earth - 0.29 - 0.37 \times 10^{22} joules - 0.23 W yr$. [In a century the depth of penetration is ~3 times more than in a decade, so heat storage in a century due to 1°C warming is ~ 0.7 W yr/m².]

¹Note that 1 W sec = 1 joule, # sec/year = $\pi \times 10^7$, area Earth = $5.1 \times 10^{18} cm^2$, 1 W yr over full Earth ~ 1.61×10^{22} joules, ocean fraction of Earth ~ 0.7, 1 calorie = 4.19 joules.

James Hansen, et al, *Earth's Energy Imbalance: Confirmation and Implications*, SCIENCE 308 (2005), p. 1431

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Permafrost Melt

Where is the permafrost?

Average latitude of permafrost boundary: 61°

(yellow circle)

(Aileen Zebrowski)

Isolated
Sporadic
Discontinuous
Continuous

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How much carbon would be released from the permafrost if the global mean temperature rose by 2 degrees Celsius?

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Budyko's Equation

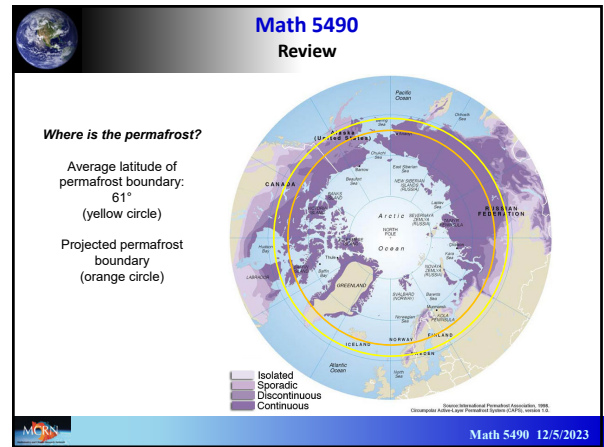
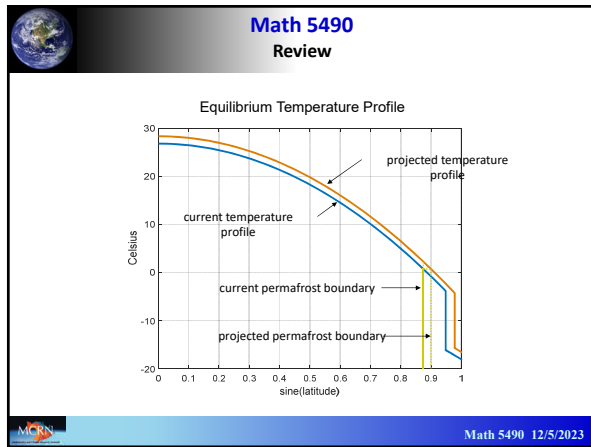
$$R \frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y)) - (A + BT) + C(\bar{T} - T)$$

Labels: surface temperature, sin(latitude), $\bar{T} = \int_0^1 T(y) dy$, heat capacity, insolation, albedo, OLR, heat transport

Symmetry assumption: $0 \leq y = \sin(\text{latitude}) \leq 1$

Chylek and Coakley's quadratic approximation:
 $s(y) \approx 1 + s_2(3y^2 - 1)$, where $s_2 = -0.241$

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How much carbon could be released from the permafrost if the global mean temperature rose by 2 degrees Celsius?

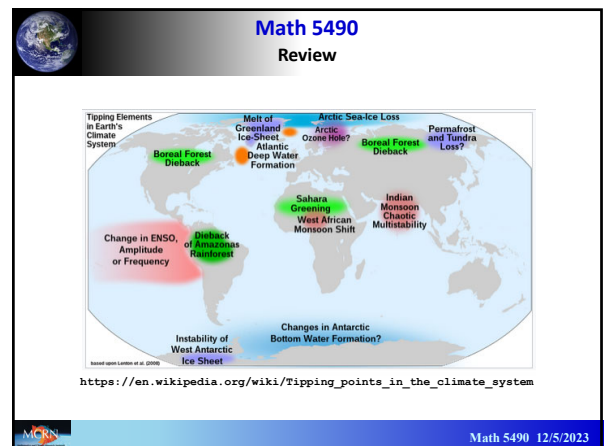
Recall that the surface area is proportional to y , the sine of the latitude.

Current permafrost boundary: $y_p = \sin(61^\circ) \approx 0.875$
Proportion of globe cover by permafrost: $1 - y_p = 0.125$
 $\Delta y \approx 0.027$

Proportion of permafrost melted: $\frac{0.027}{0.125} = 0.216$
Amount of carbon released: $0.216 \times 1400 = 302 \text{ GtC}$
Total fossil fuel emissions since 1751: 580 GtC

To hold the GMT at 2°C , we will have to withdraw 300 GtC from the atmosphere as the permafrost melts.

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https://en.wikipedia.org/wiki/Tipping_points_in_the_climate_system

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Can We Predict the Future?

If we know the state of a system now, do we know its state in the future?

For models based on differential equations, the answer is 'yes'.

$$\frac{dx}{dt} = f(x), \quad x \in \mathbb{R}^n, \quad x = x_0 \text{ when } t = 0$$

If f is sufficiently smooth (e.g., continuously differentiable) then there is a unique solution of the differential equations satisfying the initial condition.

Interpretation:

If we know the state of the system now, we can compute its state in the future.

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Math 5490
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Can We Predict the Future?

$$\frac{dx}{dt} = f(x), \quad x \in \mathbb{R}^n, \quad x = x_0 \text{ when } t = 0$$

If we know the state of the system now, we can compute its state in the future.

Yes, but how accurately?

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Example

$$\varphi(x) = 4x(1-x)$$

Lyapunov multiplier

$$\mu(x_0, n) = \left(\prod_{k=0}^{n-1} |\varphi'(x_k)| \right)^{1/n}$$

Evidence indicates that

$$\mu(x_0) = \lim_{n \rightarrow \infty} \mu(x_0, n) = 2$$

Interpretation

On average, the error multiplies by a factor of 2 at each step. After n steps, an error of ϵ becomes about $2^n \epsilon$.

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Summary

orbit: $x_{n+1} = f(x_n) = f^{n+1}(x_0), \quad x_0 = x(0)$

Lyapunov multiplier

$$\mu(x_0) = \lim_{n \rightarrow \infty} \mu(x_0, n) = \lim_{n \rightarrow \infty} \left(\prod_{k=0}^{n-1} |f'(x_k)| \right)^{1/n}$$

Lyapunov exponent


$$\lambda(x_0) = \lim_{n \rightarrow \infty} \lambda(x_0, n) = \lim_{n \rightarrow \infty} \frac{1}{n} \sum_{k=0}^{n-1} \log |f'(x_k)|$$

Interpretation

If the Lyapunov exponent is greater than zero ($\lambda(x_0) > 0$) or, equivalently, the Lyapunov multiplier is greater than one ($\mu(x_0) > 1$), then nearby orbits diverge exponentially.


If the Lyapunov exponent is less than zero ($\lambda(x_0) < 0$) or, equivalently, the Lyapunov multiplier is less than one ($\mu(x_0) < 1$), then nearby orbits converge exponentially.

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Positive Lyapunov exponents are associated with unpredictability.
Errors propagate exponentially.
Orbits behave randomly.
Simple deterministic systems can have random behavior.

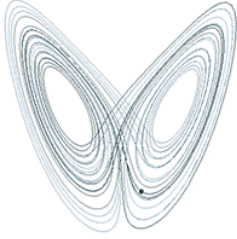


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The Lorenz Attractor


$$\begin{aligned}\dot{x} &= -\sigma x + \sigma y \\ \dot{y} &= \rho x - y - xz \\ \dot{z} &= -\beta z + xy\end{aligned}$$

The Lorenz attractor has a positive Lyapunov exponent.
i.e., nearby solutions diverge exponentially.

http://en.wikipedia.org/wiki/Lorenz_system



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